## Research note

# A decomposition theorem for the transformation of local physical frames with geodetic interest 

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Received 1977 March 7

Summary. In the present note we express the 'natural' covariant and contravariant moving frames in the actual gravity field of the Earth in terms of a 'normal' and a 'disturbing' part, the latter being expressed by the components of the deflection of the vertical and of the disturbing potential at $P$.

An elementary displacement of a point $P$ in the actual gravity field of the Earth is expressed by the differential relation
$d P=\left(\frac{\partial P}{\partial x^{*}}\right)^{T} d x^{*}=\mathrm{v}^{* T} d x^{*}$
where
$\left(\frac{\partial P}{\partial x^{*}}\right)=\left(\frac{\partial P}{\partial x^{* 1}} \frac{\partial P}{\partial x^{* 2}} \frac{\partial P}{\partial x^{* 3}}\right)^{T}=v^{*}=\left(v_{1}^{*} v_{2}^{*} v_{3}^{*}\right)^{T} ; \quad x^{*}=\left(x^{* 1} x^{* 2} x^{* 3}\right)^{T}$,
$x^{*}$ being the natural coordinates (see Marussi 1949), $x^{* 1}=\Phi$ (astronomic latitude), $x^{* 2}=\Lambda$ (astronomic longitude), $x^{* 3}=W$ (actual geopotential) and $\mathbf{v}^{*}$ the local covariant natural frame giving, in general, the displacement reference frame; $\mathbf{v}_{1}^{*}$ directed towards the actual meridian, $\mathbf{v}_{2}^{*}$ towards the actual parallel and $\mathbf{v}_{3}^{*}$ towards the actual isozenithal. Here the term 'actual' has the meaning of the term 'geodetic' used by Marussi 1952.

Decomposing $\mathrm{v}^{*}$ and $x^{*}$ in a normal and a disturbing part
$\mathbf{v}^{*}=\mathbf{v}+\delta \mathbf{v}$
$x^{*}=x+\delta x$
where, and from now on, the terms without asterisks are the normal part, referred to a uniquely defined model field, e.g. the Somigliana-Pizzetti ellipsoidal type of normal field, and the terms with $\delta$ the disturbing part; (1) is written
$d P=(\mathrm{v}+\delta \mathrm{v})^{T}(d x+d \delta x)$

[^0]or
$d P=\mathbf{v}^{T} d x+\left(\mathbf{v}^{T} d \delta x+\delta \mathbf{v}^{T} d x+\delta \mathbf{v}^{T} d \delta x\right)$.
Obviously $x=\left(x^{1} x^{2} x^{3}\right)^{T}$ are the normal coordinates at $P, x^{1}=\phi, x^{2}=\lambda, x^{3}=U$ and $\delta x=\left(\delta x^{1} \delta x^{2} \delta x^{3}\right)^{T}$ are the coordinate-disturbances, where $\delta x^{1}=\xi, \delta x^{2}=\epsilon=\eta / \cos \phi$, (the deflections of the vertical) and $\delta x^{3}=T$, (the disturbing potential) at $P$.

Denoting by $u^{*}=\left(u_{1}^{*} u_{2}^{*} u_{3}^{*}\right)^{T}$ the local contravariant natural frame, which describes gradient vectors, we have by definition, (see, e.g. Marussi 1976, p. 7)
$\mathbf{u}^{*}=\operatorname{grad} x^{*}$
with
$\mathrm{u}^{*} \mathrm{v}^{* T}=I$
where $I$ the $3 \times 3$ unit matrix.
Decomposing again $u^{*}$ and $x^{*}$, within (5), in a normal and a disturbing part, we obtain $\mathbf{u}^{*}=\mathbf{u}+\delta \mathbf{u}$
and introducing (7) and (3) into (5)
$\mathbf{u}^{*}=\operatorname{grad} \boldsymbol{x}+\operatorname{grad} \delta \boldsymbol{x}$
where obviously
$\mathbf{u}=\operatorname{grad} \boldsymbol{x}$
and
$\delta u=\operatorname{grad} \delta x$.
Introducing the astronomic orthonormal base $\mathbf{i}^{*}=\left(\mathbf{i}_{1}^{*} \mathbf{i}_{2}^{\mathbf{*}} \mathbf{i}_{3}^{*}\right)^{T}$ at $P$, where $\mathbf{i}_{1}^{*}$ is directed towards the actual north, $\mathrm{i}_{2}^{*}$ towards the actual east and $\mathrm{i}_{3}^{*}$ towards the actual zenith, the triad being clockwise, we could express $u$ by means of $i^{*}$ through a $3 \times 3$ matrix ( $A+\delta A$ )
$\mathbf{u}=(A+\delta A) \mathbf{i}^{*}$
where $A$ the normal part of the matrix and $\delta A$ the disturbing part. The normal part $A$, concerning a Somigliana-Pizzetti type of normal field, is given by Marussi (1952, p. 83, equation V-27, or: 1976, p. 30, equation 4.3) referred to the normal orthonormal base $\mathbf{i}=\left(\mathbf{i}_{1} \mathbf{i}_{2} \mathbf{i}_{3}\right)^{T}$ at $P$, where $\mathbf{i}_{1}$ is directed towards the normal north, $\mathbf{i}_{2}$ towards the normal east and $i_{3}$ towards the normal zenith, the triad being clockwise. The normal part $A$ could be also derived for any other model field, for instance the SAO satellite gravity model. The disturbing part $\delta A$ relates the $\mathbf{i}$ and the $\mathbf{i}^{*}$ frames through rotations involving the deflection components $\delta x^{1}$ and $\delta x^{2}$ and their partial derivatives, (see, e.g. Livieratos 1976, appendix 1.)

Furthermore we can express the gradients of the scalar quantities $\delta x$ by means of $u$ through a $3 \times 3$ Jacobian matrix $B$
$\operatorname{grad} \delta x=B u$
(see, e.g. Finzi \& Pastori 1971, p. 21, section 4). The Jacobian matrix B involves the partial derivatives of $\delta x$ with respect to the normal coordinates $x$. Introducing (11) into (12) we obtain
$\operatorname{grad} \delta x=B(A+\delta A) \mathrm{i}^{*}$.
Replacing (11) and (13) into (7) we can write, with the help of (8)
$\mathbf{u}^{*}=(A+\delta A) \mathrm{i}^{*}+B(A+\delta A) \mathrm{i}^{*}$
or
$\mathbf{u}^{*}=(A+B A+\delta A+B \delta A) \mathrm{i}^{*}$.
Using the condition (6), and having by definition
$\mathrm{i}^{*} \mathrm{i}^{* T}=I$
we can write
$\mathbf{v}^{*}=\left[(A+B A+\delta A+B \delta A)^{T}\right]^{-1} \mathbf{i}^{*}$
if and only if $\operatorname{det}(A+B A+\delta A+B \delta A) \neq 0$.
The transformation matrices in (15) and (17) relate the covariant and the contravariant local frames with respect to the astronomic orthonormal base $\mathbf{i}^{*}$ at $P$, and depend only upon three simple $3 \times 3$ matrices $A, B$ and $\delta A$, the matrix $A$ being the normal part and the matrix $(B A+\delta A+B \delta A)$ the disturbing part.

Having already established the local covariant vector frame $\mathbf{v}^{*}$, it is easy to derive the metric tensor $M^{*}$ in the actual gravity field, defined by the relation
$M^{*}=\mathrm{v}^{*} \mathrm{v}^{* T}$.
Introducing (17) into (18) we obtain
$M^{*}=\left[(A+B A+\delta A+B \delta A)^{T}\right]^{-1} \mathbf{i}^{*} \mathbf{i}^{* T}(A+B A+\delta A+B \delta A)^{-1}$
and recalling (16)
$M^{*}=\left[(A+B A+\delta A+B \delta A)(A+B A+\delta A+B \delta A)^{T}\right]^{-1}$.
Performing matrix algebra, we obtain the metric tensor decomposed into a normal and a disturbing part

$$
\begin{align*}
M^{*}= & \left(A A^{T}+\right. \\
& +A A^{T} B^{T}+A \delta A^{T}+B A A^{T}+\delta A A^{T}+ \\
& +A \delta A^{T} B^{T}+B A \delta A^{T}+\delta A A^{T} B^{T}+\delta A \delta A^{T}+B \delta A A^{T}+B A A^{T} B^{T}+  \tag{21}\\
& +B A \delta A^{T} B^{T}+\delta A \delta A^{T} B^{T}+B \delta A A^{T} B^{T}+B \delta A \delta A^{T}+ \\
& \left.+B \delta A \delta A^{T} B^{T}\right)^{-1} .
\end{align*}
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The reciprocal components of the metric could be as well derived from the definition $u^{*} u^{* T}$.

## Acknowledgments

Thanks are due to Professor A. Marussi, Trieste, for his discussions on the subject. Suggestions given by Professor E. Grafarend, Munich, are kindly acknowledged.

## References

Finzi, B. \& Pastori, M., 1971. Calcolo tensoriale e applicazioni, II, ed., Zanichelli, Bologna.
Livieratos, E., 1976. On the geodetic singularity problem, Manuscripta Geod., No. 4, W. Berlin.
Marussi, A., 1949. Fondements de géométrie differentielle absolue du champ potentiale terrestre, Bull. Geod. No. 14, Paris.
Marussi, A., 1952. Intrinsic geodesy, Map. Chart. Res. Lab., OSU Techn. Rep. 159, Columbus.
Marussi, A., 1976. Intrinsic geodesy, Boll. Geod. e Sc. Aff., No. 1, Firenze.


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